Comments to Bonfante et al’s manuscript – v3 – 23 April 2015

General comments
Again I wish to emphasize that is a very interesting and innovative approach combining soil (physical) information, the use of a deterministic model and the chain vine-grape-wine data which deserves publication in the SOIL journal. Some improvements were brought to this version of the text. The title was modified as requested. However, there still remain some additional modifications not addressed according to reviewer’s comments, and which are likely to improve the final version of this manuscript prior to publication.

In particular, it is still unclear to the reader whether each of the 27 plants monitored in each HZ is unchanged over the 3 monitoring years (as suggested in section 2.8) or whether these 27 plants are likely to change and were randomly sampled at each measurement date (as suggested in section 2.7). Please clarify this. In the first case, the 54 vines locations could be shown on the map, which is impossible for the second case.

The plants were randomly sampled, now in the text this information is clearly reported.

How was the set of 10 plants for Leaf water potential measurement selected? Was it included amongst the 27 monitored vines? Was it a predawn or midday Leaf water potential measurement, or several per-day measurements?

The 10 plants monitored were randomly chosen between the 27 plants monitored in each fHZ for the crop monitoring. The leaf water potential was midday. Now this is clarified in the text.

As already suggested by reviewers, section 3.1 lacks an additional topographical map showing sampled profiles, and/or pits and augers locations which shall facilitate the reading of this text.

The topographical map showing sampled profiles, and/or pits and augers locations has been added (fig. 3a)

A photograph or sketch of typical soil profiles would be welcome.

Photos of both soil profiles have been added.

(In absence of location map, the reader may have to read several times as the location of the Calcisol on the upslope instead of downslope is rather unexpected...questioning the origin of its Bk horizon; is the underlaying clay a marl rather than a clay? ... Where does the Cambisol colluvium originate from? It is surprising how low its carbonate content is despite its downslope position regarding to the upslope clay (marl?) and Bk horizon...such questions are of course beyond the scope of this article).

Yes we agree that many of these interesting questions posed are far beyond this article.

Anyway for the sake of clarifying the issue ... our Calcisol develop on clayey sediments rich in carbonates. This is rather common over hilly landscapes in marl-like Apennines. Possibly this is also related to ancient geomorphic erosional processes. Cambisol colluvium seem to have a (partly) volcanic component (no idea about age) since there are some pumices in coarse fragments. But NO andic properties occur.

Conversely to the author’s statement, both soils do not have quite the same texture. In lower horizons (Bw2 and Bw3), the Cambisol has higher clay content compared to Calcisol and this results in clay texture...This may impact vine response depending on root depth. In which horizons were the roots observed and considered “efficient”, particularly in the SWAP modelling?

Our statement refer to the soil classification, which is based on the weighted average along the soil profile of particle soil size (clay, silt and sand). Generally, in agricultural sectors as the viticulture, soils are often classified on this weighted basis. Our approach is obviously oriented to emphasize the soil physical
differences and behaviors due to presence of different soil horizons (thickness and vertical depth and position inside of profile). These information are present into the physically based model application to solve the soil water balance (E.g. SWAP).

About the roots depth, obviously the soil physical characteristics influences the roots vertical distribution and density. In our case, the increments of clay in the BW2 and BW3 horizon doesn’t produce a massive soil structure but help to produce porosity, moreover, in these two horizons the organic matter is (1.6%) comparable with those present in the upper layers of Calcisol.

During the soil profiles descriptions the roots distribution was recognized. In particular, the Calcisol roots distributions was within the first 65 cm of soil, few roots were present beyond this depth. This behavior was, in our opinion, due to the presences of calcium concentrations of Bk horizon (see the white spots in Bk horizon in the figure 4) which may have influenced roots distribution.

In the cambisol there were none physical or chemical constrains, but the roots distribution was within the first 85 cm of soil depths. In our opinion, on the base also of our campaign of soil monitoring, this could be due to the water availability, that it does not force the plant to go deeper with roots.

In SWAP modelling we have described the roots vertically density and depth in both soils in according with what we have seen in field.

Discussion section: Potential limitations of this approach? In particular, it should be emphasized that the within-HZ spatial variability was not accounted for but may impact the uncertainty of predictions (Table 2).

We have added a sentence in the manuscript (line 614 new version): “Moreover, our approach could be improved towards the precision farming considering the soil spatial variability inside of each fHZs in the modelling application, thereby providing also an information about the uncertainty of model predictions. But, the used of latter approach is strictly limited by the availability of soil information spatially distributed needed to apply the simulation model (verticals and horizontals information, e.g. soil horizons depth and thickness and their hydraulic properties).”

In our case study, this approach was not used also because the scale of approach was different and because our work does not aim to deliver towards precision farming approaches.

Did the authors test SWAP modeling at each of the individual profiles within HZs and if so, which variability did they obtain?
The SWAP model was applied to each representative soil profile identified for each HZs. They represent a real soil profile and not a constructed one.

Are there other variability factors not addressed by the SWAP model?

SWAP model is unidimensional model addressed to solve the soil water balance considering the soil, plant and atmosphere system. Obviously, as all simulation model, it represent a simplification of reality, then some aspects variable in the space affecting the plant responses could be not taken into account. However, in our approach, the physically description of soil water balance and the use of simple crop development are able to give to us the information that we need (CWSI). Moreover, the use of simple crop model is very important in a procedure addressed to the viticultural zoning planning. The use of more sophisticated model on crop growth needs the knowledge of a specific value for many variables used to simulate the crop growth. This without a local and cultivar specific calibration increase the model prediction uncertainty (most of the relation applied in the crop growth are empirical and not physically bases). Finally in our case study a dataset for the Aglianico grapevine is not present in literature (perhaps also for the most important grapevines in the world).
Tables and figures

Figure 3: ECa values should be expressed into data ranges instead of single values. How where these values thresholded? Same questions for Figure 5.

The ECa values of figure 3 and 5 are now reported in terms of range. The thresholds were chosen in accordance to the values measured in proximity of the soil profiles and augers realized during the soil survey.

Table 1: Were these properties averaged for the profiles categorized in each HZ? Or do they correspond to a single representative profile? Or an "ideal" modal profile synthetized from all observations?

In any case, it would be worth showing the location of each of these representative profiles in the maps as already suggested by reviewers.

The values reported in the table 1 correspond to a single representative profile. The map with location of soil profile was added in the manuscript.

The properties Q0, a, l should be defined; "a" (isn't alpha?) as property is likely to be confused with "a" "absence of rock fragments".

Please write particle size fraction instead of texture.

OK done.

Figure 2 and Figure 4 and section 3.2: it would be worth showing the same plots computed for or over the three monitored years: is the monitored period representative of the 2003-2013 series?

The same representative soil profiles were computed in both figures. The monitored period (2011-2013) surely is representative of recent climate (2003-2013) because, the year 2012 was very dry year (about 200 mm of rainfall during the cropping season), the 2011 and 2013 can be considered as a normal years (about 300 mm of rainfall during the cropping season) because the average rainfall during the recent climate period was of 320 mm (+-112 mm).

Table 2 and section 3.2: the SWAP simulation results should also be given for the 3 monitoring years...to be comparable to plant and harvest monitoring results.

The use of model in this approach is not oriented to evaluate the relation between the monitored years and model results, but to validate the use of simulation model application in the viticultural zoning at farm scale, in order to identify the functional homogeneous zones fHZs. What we want to stress are the potentiality of the simulation model application (defined by CWSI) in the viticultural zoning planning. Thus, we are afraid that the introduction in the discussion of the relation between the model results in the three years of monitoring will create confusion in the reader and the focus of paper will be lost. Thus, for those reasons in our discussion we would like to avoid to move away from our principal goal, which was to show the potential of the use of simulation model in viticultural zoning planning.

In detail:
Line 134, page 3: 368 m a.s.l. instead of 368 a.s.l. Done
Line 142, page 3: please specify in what this “green manuring” consisted (doses/ha/y, nature?). Each year, the farmer do only the legume green manure. Done
Line 228, page 6, Van Leeuwen instead of Van Leewen Done
Line 282, page 7, S instead of Sis
Lines 402-404: how were the 10 augers and 6 pits locations localized? Figure added
Line 410, page 10, how high is the active lime content of the Calcisol? Are the vines planted on Calcisols suffering chlorosis (Lines 507-512: is the lower chlorophyll content for Calcisol explained by water stress only)? The 1103 Paulsen rootstock is known to have a rather limited resistance to chlorosis.
The 1103 Paulsen was selected in southern Italy for its strong drought tolerance and its ability to grow well on lime-based soils. In literature is classified as: rootstock with high tolerance to drought (Lavrencic et al., 2007) or resistent to drought (Carbonneau, 1985) and with medium tolerance to chlorosis (Chauvet & Reynier, 1979). But in this last case, if we consider the “chlorotic power index” (CPI) (Juste & Pouget, 1972.
In: Huglin & Schneider, 1998) and the degrees of chlorosis intensity in relation to different values of CPI (Lupascu et al., 2009), seems that the 1103P rootstock have an high intensity of chlorosis (classified as value CPI=30, in Huglin & Schneider, 1998). But, in our case study, none evidences of chlorosis were recognized by the farmer from the planting of vineyards (2000) to today in both soils. This could be due to the Aglianico cultivar. The local farmers claim that the chlorosis with 1103P manifests more on white vines as Falanghina cv. and not on red vine as Aglianico. Only the vigor effects induced by the presence of water availability in the cambisol was recognized in field. From Christensen, L.P. 2003. Rootstock Selection (PDF). Pages 12-15 in: Wine Grape Varieties in California. University of California Agricultural and Natural Resources Publication 3419, Oakland, CA. http://iv.ucdavis.edu/files/24347.pdf

Moreover, the differences in chlorophyll content between the plants in the CAD and CAM HZs were not symptom of chlorosis (the leaves were a less intense green color, diffused and not speckled). The differences could be attributed to soil water stress because the crop management was the same between the two soils.

Line 414, page 10, EC or ECa? Solved
Line 415, page 10, “the texture is clay loam in both soils” this does not hold true for lower horizons. Please moderate the sentence. Specify the main rooting depth in each case. See the above comments
May the deep clay texture explain the highest vigour observed for Cambisol compared to Calcisol, all the more than the 1103 Paulsen rootstock might induce vigour.

As reported above, the 1103P effects on plant vigor induced by the presence of water availability in the cambisol has been recognized. The question is that is not the clay per se to produce the vigor effect, but the presence of hydraulic properties (as retention curve) able to produce a better condition for the water availability. The presence of clay can improve the soil structure (porosity) if works as colloid, obviously if the pedogenesis processes are began in the soil. But there is not a directly correlation between soil structure and clay content, it depends from the soil formation factors (Climate, Organisms, Relief, Parent material and Time, CRORPT - Jenny 1941). If the pedogenesis processes are not started in the soil, the presence of high clay content doesn’t mean that there are better conditions for water availability for the plant (no soil particle aggregation is began). In fact, very often, in the deep soil horizons, the high clay content correspond a massive soil structure and not in a soil porosity increase. For those reasons is good to have the hydraulic measurement measured in field and not only by means of PTF.

Line 418, page 10, K₀ instead of k₀; what is “I”? 
The parameter I (-) is an empirical pore tortuosity/connectivity parameter applied to describe the hydraulic conductivity of unsaturated soils (e.g. in the van Genuchten-Mualem model, van Genuchten, 1980).
Functional homogeneous zones (fHZs) in viticultural zoning procedure: an Italian case study on Aglianico vine.

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Abstract

This paper aims to test a new physically oriented approach to viticulture zoning at farm scale, strongly rooted in hydopedology and aiming to achieve a better use of environmental features with respect to plant requirements and wine production. The physics of our approach is defined by the use of soil-plant-atmosphere simulation models, applying physically-based equations to describe the soil hydrological processes and solve soil-plant water status.

This study (ZOVISA project) was conducted on a farm devoted to high quality wines production (Aglianico DOC), located in South Italy (Campania region, Mirabella Eclano-AV). The soil spatial distribution was obtained after standard soil survey informed by geophysical survey. Two Homogenous Zones (HZs) were identified; in each one a physically based model was applied to solve the soil water balance and estimate the soil functional behaviour (crop water stress index, CWSI) defining the functional Homogeneous Zones (fHZs). For the second process, experimental plots were established and monitored for investigating soil-plant water status, crop development (biometric and physiological parameters) and daily climate variables (temperature, solar radiation, rainfall, wind).

The effects of crop water status on crop response over must and wine quality were then evaluated in the fHZs. This was performed by comparing crop water stress with (i) crop physiological measurement (leaf gas exchange, chlorophyll-a fluorescence, leaf water potential, chlorophyll content, LAI measurement), (ii) grape bunches measurements (berry weight, sugar content, titratable acidity, etc.) and (iii) wine quality (aromatic response). This experiment proved the usefulness of the physical based approach, also in the case of mapping viticulture microzoning.

Key Words: Terroir, SWAP model, LWP
1 Introduction

Concepts such as terroir and viticulture zoning are becoming increasingly more important for planning and managing vineyards aiming at high quality wine (Gladstones and Smart, 1997; Carey, 2001; Vaudour, 2003). Basically their practical implementation (Deloire et al., 2005; Fregoni, 1988) aims to classify the landscape (mainly climate and soil), studying its interaction with vineyard and wine quality. Mapping of terroir and viticulture zoning have been developed at all scales, especially since the 1990s following the widespread use of geomatics (Girard and Girard, 2003). The methodology (even if not unique) indeed gave many positive results, but also showed some important limitations mainly related to its strongly empirical base. In other words the terroir is a sort of “black box” in which the quantitative linkage between climate–soil–plant and wine is empirically or statistically described (e.g. Brousset et al. 2010) and not analyzed in its mechanics (Bonfante et al. 2011).

Recently some changes have been made and the spatial analysis of terroirs has improved, incorporating some key features known to strongly affect wine quality. Among them are solar radiation and bioclimatic indexes (Failla et al., 2004; Vaudour, 2003) and also morphometric data and multitemporal remotely sensed images (Vaudour et al. 2010). Moreover, Bonfante et al. (2011) demonstrated that terroir analysis – applied at a district scale (mesoscale sensu Shaw & Shaw 2005) – can become more profitable by combining high quality GIS (as bioclimatic indexes) with water balance simulation modelling for addressing the key and very complex issue of soil-plant water stress. This is very important because even if it is well known that water stress strongly affects grape quality, its spatial description can be a very difficult issue.

Despite this result, it is still very questionable whether a similar approach can be usefully applied at a more local – less aggregated – spatial scale where ecophysiological functioning and land management play a key role. Moreover, this detailed scale is very useful because it makes possible to evaluate the functional relationships between viticulture zones, plant-soil water stress, vineyard status and grape wine quality.

To address this issue here we refer to “micro-zoning” in coherence with the term “micro-scale” used by Vaudour & Shaw (2005) for terroir zoning.

In this perspective, the aim of this paper is to prove that physically based approaches can be usefully employed also at very detailed scales such as for viticulture microzoning, in order to effectively separate different viticulture zones (Functional Homogeneous Zones, fHZ) on the basis of their potential functionality (e.g. potential water stress) and by doing so better orient viticulture management.

This was done on an experimental site (2.3 ha over a homogeneous hilly slope) characterized by large soil variation under the very same climatic conditions.

2 Materials and methods

2.1 Study Area

The study area is located in a hilly environment of southern Italy (Mirabella Eclano - AV, Campania region: Lat. 41.047808°, Lon. 14.991684°, elev. 368 m.s.l.), in a farm oriented to the production of high quality wines, namely Aglianico cultivar (controlled designation of origin –DOC /AOC), standard clone population planted in the year 2000 on 1103 Paulsen rootstocks (espaler system, cordon spur pruning, 5000 units per hectare) placed along a slope of length 90 m with an 11% gradient. Legume green manure, management is applied.

The long-term (2003-2013) mean daily temperature annually of the study area was 14.7 (±0.9) °C, while the mean annual rainfall was 802 (±129) mm (data from the Regional weather station of Mirabella Eclano – AV- at 1 km of study area)
Climate monitoring within the farm during the 2011 to 2013 vintage showed that during the cropping season (April to early October) the mean daily temperature was 20.9 (±1.2) °C, while the precipitation was very variable during the three vintages, ranging from 285 to 200 mm.

### 2.2 Method used for mapping Homogeneous Zones

The viticulture microzoning procedure used is rooted in procedures already applied at various scales. This includes standard soil mapping and geometric spatial analysis of solar radiation, bioclimatic indexes, morphometric data and remotely sensed images (Failla et al. 2004; Vaudour 2001; Vaudour et al. 2010). Moreover, the procedure included geophysical mapping which has proved to be a very useful tool for soil spatial analysis also in precision viticulture (Andrenelli et al. 2010; Priori et al. 2010; Priori et al. 2012).

More specifically, the employed procedure is given in the flow diagram in fig. 1:

- **Step 1**: Identification of Homogeneous Zones (HZs) obtained from standard soil mapping (landscape units, soil profiles, minipits, etc.) at a detailed scale supported by geophysical survey. These HZs were also statistically described in terms of their DTM and DSM derived parameters.

- **Step 2**: Evaluation of the hydrological indicator of crop water status (potential crop water stress index - CWSI) by applying physically based simulation modelling to the representative soils of the previously defined HZs, and definition of functional Homogeneous Zones (fHZs).

- **Step 3**: Vine/must/wine monitoring over the fHZs.

The realization of step 1 gives an advanced but “static” description of landscape useful for making a standard Land Evaluation (qualitative and empirical approach) to evaluate how suitable the land is for the vine. The innovation is enclosed in step 2, where a key component of the functional behavior of soils to vine responses is described dynamically by means of a physically based approach (Bonfante et al. 2011). This last step makes it possible to discriminate soil behavior through a hydrological indicator of plant water status (Crop Water Stress Index - CWSI) and to identify the functional Homogeneous Zones (fHZs) from the HZs. The term “functional” is employed in order to strengthen the soil-plant-climate functionality. Finally, step 3 allows evaluation of plant behavior within any fHZs and testing of the occurring CWSI.

**Fig. 1.** Flow diagram illustrating the proposed approach applied to the case study of Quintodecimo farm.
2.3 Pedological survey and soil measurements

A combined geophysical-pedological approach was used to derive the map of the Homogeneous Zones (HZs) in the study area. These HZs were obtained after a detailed soil survey that was adapted to the specific need of this research. A preliminary map of the most important soil forming factors in the area was obtained by combining a geomorphological analysis of LiDAR-derived DEM with local geological information. This led to the identification of the main soil-landscape units. The actual soil survey (soil profiles, minipit and augering) was supported by geophysical survey techniques. These techniques, based on non-invasive tools (generally applied to environmental studies, e.g. for geological prospecting), were used as very quick survey systems that gave initial information on the general spatial variability of the soils that were investigated, better planning of the field investigation in the pedological survey, and improvement in the soil map resolution emphasizing the spatial soil micro-variability (traditional soil surveys and soil analysis are usually time-consuming and expensive, especially for high resolution maps). Geoelectrical soil mapping has become widely accepted and considered as a successful geophysical method that provides the spatial distribution of relevant agronomic information for precision farming (Lück et al., 2009).

Methods based on electrical properties are particularly promising as support to pedological surveys because important soil physical properties are strongly correlated to electrical conductivity which, changing in space, can represent spatial soil distribution. Geophysical methods offer a valuable means for obtaining subsidiary data in an efficient way, and have been widely applied in soil sciences for a considerable period of time (Samouelian et al., 2005). In this study the apparent soil electric conductivity (ECa), was carried out by Electro-Magnetic Induction (EMI) sensors, which represents a very useful tool for identifying soil map units and soil properties in respect of clay content (Morari et al., 2009), soil depth (Saey et al., 2009), water content (Davies, 2004; Cousin et al., 2009; Lück et al., 2009; Tromp-van Meerveld and McDonnell, 2009) and water salinity (Doolittle et al., 2001).

However soils, like every other geological material, are not uniform, consequently what is specifically measured is an apparent electrical conductivity (ECa), which can be defined as the actual conductivity of a rock homogeneous and isotropic equivalent to a real heterogeneous and anisotropic media.

The instrument used for surface mapping the electric conductivity was the EM38-DD (Geonics Ltd., Ontario, Canada) used in both VDM (vertical dipole mode) and HDM (Horizontal dipole mode). The sensor of instrument was calibrated to minimize the errors before the survey, which was performed in July 2011 during grape ripening. The instrument was placed on a PVC sledge and pulled by a tractor along the inter-rows, at a distance of about 5 meters to avoid interference phenomena. The use of the sledge makes it possible to keep the instrument at a constant distance from the soil, making data acquisition easier and more accurate.

The data were recorded on a GPS-supplied data-logger with European Geostationary Navigation Overlay Service (EGNOS)–Wide-Area Augmentation System (WAAS) correction (accuracy ± 3 m), which made it possible to georeferenced and map the measured property. The instrument was set to acquire one measurement per second.

Data post-processing was performed by ordinary kriging with 1 m resolution. The final result of the EM38-DD survey was therefore a regular grid of data points including ECa for two depths (1.6 m for VDM and 0.76 m for HDM). These horizontal (HDM) and vertical (VDM) ECa maps were used as baseline data for a pedological survey based on soil augerings and soil profiles descriptions. This was done similarly with the classic (overlying procedures) soil survey of other thematic layers (geology, geomorphology, etc.).

The soil profiles were described according to FAO (2006), Chemical analyses were performed according to the official methods of the Italian Ministry of Agriculture and Forestry REF. The grain size distribution (GSD) was determined by a laser granulometer (MalvernMastersizer 2000).

Undisturbed soil samples (volume ≥ 750 cm³) were collected from each soil horizon and hydraulic properties were determined in the laboratory to simulate the hydrological conditions of the soil by means of an agro-hydrological model, which is illustrated in section 2.5.

Soil samples were saturated from the bottom and the saturated hydraulic conductivity was measured by a permeameter (Reynold et al., 2002). After sealing the bottom surface to set a zero flux, measurements were then taken during drying: at approximately pre-set time intervals, the weight of the whole sample and the pressure head at three different depths (by means of tensiometers) were determined. An iterative procedure was applied for estimating the water retention curve from these measurements. The instantaneous profile method was used to determine the unsaturated hydraulic conductivity. Moreover, some points at a lower water content of the dry branch
of the water retention curve were determined by a dew-point system (WP4 dew-point potentiometer, Decagon devices, Inc.). Details on the tests and overall calculation procedures were described by Basile et al. (2012) and Bonfante et al. (2010).

2.4 The hydrological indicator: Crop Water Stress Index (CWSI)

The effects of water stress on wine quality, appearance, flavour, taste and aroma have been clearly highlighted by different authors: Matthews et al. (1990) Van Leeuwen et al., 2009, Chapman et al., 2005; Acevedo-Opazo et al., 2010, Intrigliolo and Castel 2011 and Romero et al., 2013, differentiating also the effects between early or late water deficit treatments. The estimate of water stress at the different phenological stages can indeed, therefore, represent an important tool in terroir classification. Different variables (e.g. air temperature, wet-bulb temperature, etc.) could be applied to develop a proper water stress index, but as reported by Kozak et al., 2006, the use of transpiration information is realistically more variable (in respect to Evapotranspiration) for defining the crop water stress. In our approach, to simulate the soil water balance we used a simulation model (SWAP) based on the Richards equation, very different from the one applied by Kozak et al. It is very robust for simulating the soil water balance and, moreover, it was previously used and tested in Italy and in the same Campania Region (Bonfante et al., 2010; 2011).

The stress index estimated from the model output is a daily crop water stress index (CWSI), defined as follows:

\[ CWSI = \left(1 - \frac{T_r}{T_p}\right) \times 100 \]  

where \( T_r \) is the daily actual water uptake and \( T_p \) is the daily potential transpiration.

The sum of the daily CWSI in the required period represents the cumulated stress CWSIcum:

\[ CWSI_{\text{cum}} = \left[ \frac{\int_{t_1}^{t_2} \left(1 - \frac{T_r}{T_p}\right) \, dt}{t_2 - t_1} \right] \times 100 \]  

The application of this index, changing the integration time (\( t_1 \) and \( t_2 \)), makes it possible to estimate plant water stress at different stages of crop growth (shoot growth, flowering, berry formation, berry ripening) (Fig.2).

Finally, this index was used to analyse the PHZs behaviour and successively define the HZs.
Fig. 2. The crop water stress index (CWSI) simulated by SWAP in the HZs CAL and CAM in the year 2011 during the cropping season. (SG= Shoot growth; FL= Flowering; BF= Berry formation; BR= Berry Ripening)

2.5 Simulation modelling

The Soil–Water–Atmosphere–Plant (SWAP) model (Kroes et al., 2008) was applied to solve the soil water balance and to calculate the CWSI for each soil identified by the soil survey. It was already used in viticulture by different authors (Ben-Asher et al., 2006; Minacapilli et al., 2009; Bonfante et al., 2011; Rallo et al., 2012).

SWAP is an integrated physically based simulation model of water, solute and heat transport in the saturated – unsaturated zone in relation to crop growth. In this study only the water flow model was used; it assumes a 1-D vertical flow process and calculates the soil water flow using the Richards equation:

\[
C(h) \frac{dh}{dt} = \frac{\partial}{\partial z} \left[ K(h) \left( \frac{\partial h}{\partial z} + 1 \right) \right] - S(h)
\]

where \(C(h) = \frac{\partial \theta}{\partial h}\) is the differential soil water capacity, \(\theta\) (cm\(^3\) cm\(^{-3}\)) is the volumetric soil water content, \(h\) (cm) is the soil water pressure head, \(t\) (d) is the time, \(z\) (cm) is the vertical coordinate taken positively upward, \(K\) (cm d\(^{-1}\)) is the hydraulic conductivity and \(S\) (cm\(^2\) cm\(^{-3}\) d\(^{-1}\)) is the water extraction rate by plant roots.

Soil water retention is described by the unimodal \(\theta(h)\) relationship proposed by van Genuchten (1980), expressed in terms of the effective saturation, \(S_e\), as follows:

\[
S_e = \left( \frac{1}{1 + (\alpha h)^n} \right)^m
\]

where \(S_e = S_e(0,0,1)\) is the residual water content and the water content at \(h=0\) respectively, and \(\alpha\) (cm\(^{-1}\)) and \(n\) and \(m\) are curve-fitting parameters.

Mualem’s expression (Mualem, 1976) is applied to calculate relative hydraulic conductivity, \(K_r\). Assuming \(m=1-1/n\), van Genuchten (1980) obtained a closed-form analytical solution to predict \(K_r\) at a specified volumetric water content:
where \( K_s \) is the hydraulic conductivity measured at \( h_0 \), and \( t \) is a parameter which accounts for the dependence of the tortuosity and partial correlation between adjacent pores. The condition at the bottom boundary can be set in several ways (e.g. pressure head, water table height, fluxes, impermeable layer, unit gradient, etc.).

The upper boundary conditions of SWAP in agricultural crops are generally described by the potential evapotranspiration \( ET_p \), irrigation and daily precipitation. The potential evapotranspiration is then partitioned into potential evaporation, \( E_v \), and potential transpiration, \( T_a \), according to the LAI evolution, following the approach of Ritchie (1972).

SWAP simulates water uptake and actual transpiration according to the model proposed by Feddes et al. (1978), where the root water uptake, \( S_r \), is described as a function of the pressure head, \( h \):

\[
S_r(h) = a(h) \cdot S_{\text{max}} = a(h) \cdot \frac{T_a}{T_{\text{max}}}
\]

where \( a \) (cm) is the thickness of the root zone and \( a(h) \) is a semi-empirical function of the pressure head \( h \), varying between 0 and 1. The shape of the function \( a(h) \) depends on four critical values of \( h \), which are related to crop type and to potential transpiration rates. The actual transpiration rate \( T_a \) (cm \( \text{d}^{-1} \)) is computed by the integration of \( S_r \) over the root layer. The root depth is specified by the user as a function of the development stage.

Model parameters and data for simulations:
- the upper boundary condition comes from the daily data of the Mirabella Eclano Regional weather station (1 km from the study area) integrated with the micrometeorological station located near the farm. Daily potential Evapotranspiration (ET\( \text{p} \)) was determined by applying the Penman-Monteith equation.
- the boundary condition was set as a unit gradient.
- crop data. The Leaf Area Index was measured in different phenological phases by a ceptometer, rooting depth was measured during the profile description, the water uptake function parameters were derived from literature (Taylor and Ashcroft, 1972).
- the hydraulic properties were parameterised by fitting a procedure of the van Genuchten-Mualem model to the experimental data (see section 2.2).

The SWAP model was previously calibrated and validated in both representative soils from the study area, on the soil water content measured at different soil depths by TDR probes (5 soil depths until 100 cm) in the years 2011 and 2012 respectively. In particular, the Root Mean Square Error (Loague and Green 1991) showed values (over the soil profile) of 0.034 (±0.03) for the calcisol and 0.032 (±0.01) for the cambisol with a correlation index \( r^2 \) of 0.75 (±0.3) and 0.90 (±0.1) respectively (indexes are a weighted average over depths along the profile (until -100 cm, rooting zone).

The RMSE values agree with those shown in a previous study (Bonfante et al., 2010). Moreover, Sheikh and van Loon (2007) reported several RMSE values obtained from calibration and validation procedures by: Heathman et al. (2003), Crescimano and Garofalo (2005), Mertens et al. (2005), Sing (2005), Wegehenkel (2005) and Sheikh and van Loon (2007). Most of these results have a range of 0.03-0.05. Finally, Eitzinger et al. (2004), comparing SWAP, CERES and WOFOST models, obtained RMSE values ranging from 0.007 to 0.07 for different soils, models and crops.

We can therefore consider the application of SWAP in both soils as being good for predicting the soil water balance.

### 2.6 GIS analysis: DTM and DSM information

High resolution Digital Surface Models (DSM) and Digital Terrain Models (DTM) of the study area were acquired respectively in April 2011, 1 m spatial resolution – DTM and July 2013, 0.30 m spatial resolution – DSM with LiDAR technologies, as part of an ongoing project in the study area. These models represent the elevation values of the ground level plus those aboveground (i.e. canopy). They were processed using specific software coupled with the GIS
environment (ArcGIS, QGIS and SAGA open source software) to support the procedures of step 1 concerning (i) the
gеomatрical analysis for the pedological survey (identification of preliminary landscape mapping units), and (ii)
to investigate the variation of the study area within the Homogeneous Zones. These high spatial resolution
acquisitions gave detailed auxiliary spatial information such as the estimate of solar radiation taking into consideration
shadows from vineyards. Specifically, continuous maps of slope, aspect and topographic wetness index (TWI) were
obtained from the DTМ and potential insolation at very high spatial resolution from the DSM. More specifically, these
last derived maps were realizing the consideration of the presence of vineyard rows as 3D objects in space, able to influence
insolation through the formation of intra-rows shadows. Subsequently, all of this information was used to characterize
the differences between the HZs identified in the vineyard.

2.7 Crop measurements
Monitoring was conducted for three years (2011 to 2013) within the HZs (identified in step 2) on the vegetative
growth of 27 plants (54 plants over 2.3 ha) until the harvest. The measurements were realized randomly on a weekly
or biweekly basis, in relation to the measured variable and the physiological crop stage.

The midday Leaf Water Potential (LWP, MPa) was assessed for each HZs on a set of 10 plants randomly chosen
between the 27 plants monitored, using a Scholander type pressure bomb (SAPS II, 3115, Soil moisture Equipment
Corp., Santa Barbara CA, U.S.A). Photosynthetic CO₂ assimilation (µmol m⁻² s⁻¹), stomatal conductance to water
evapour (mol m⁻² s⁻¹) and effective quantum yield of PSII photochemistry (ΦPSII) in light-adapted leaves were measured by means of a portable photosynthesis system (LI-6400-40, LiCor, Lincoln, NE, U.S.A.). The light source was set at a saturating Photosynthetic Photon flux density of 1800 µmol m⁻² s⁻¹ while the external CO₂ source was set at 370 µmol mol⁻¹. The instrument software calculated the various gas-exchange parameters on the basis of the von
Caemmerer and Farquhar (1981) model, and ΦPSII according to Genty et al. (1989). The chlorophyll content of the
leaves was optically estimated as a relative index (CCL) using a handheld meter (CCM200, Chlorophyll content meter
Agpée Instruments, Inc., Logan, UT.) as the ratio of the fractional transmittances at 653 and 931 nm.
A linear Accupar LP-80 PAR-LAI ceptometer (Decagon Device Inc., Pullman, WA, USA) was used to measure light
interception by the vineyard and to estimate the leaf area index (LAI). The ceptometer had 80 photosynthetic photon
flux density (PPFD) sensors spaced at 1 cm intervals, and it was programmed to average readings of 10 sensors at a
time before logging data. The PPFD transmitted through the canopy (PPFDt) was measured at 0.25 cm above soil
surface over a grid of 0.1 x 0.1 cm² across an area of length 2 m and with 2 m between the rows. The measurements
were carried out in 3-4 replicates in both CAL and CAM sites, while the measurements taken in a clear area near the
two sites were taken as the PPFD incident over the canopy (PPFDI). Intercepted light (PPFDint) was calculated as the
difference between incident and transmitted PPFD, whereas the fractional light interception (fI) was calculated as the
ratio between PPFDInt and PPFDI. Statistically significant differences between the means of the analysed variables for
the two sites were evaluated using the Student’s t-test. A null hypothesis was rejected at P ≤ 0.05.

2.8 Must/wine characteristics
In addition to the crop measurements, the must and wine characteristics were monitored within the HZs (identified in
step 2) on 27 plants for three years (2011 to 2013). In particular, of the 27 plants monitored, 12 were used to collect
the grapes at harvest and 15 for sampling scalar grapes (randomly-sampled at each measurement date).

The Standard chemical analyses and spectrophotometric measurements of must and wine were carried out as follows:
Standard chemical analyses (soluble solids, total acidity, pH, total polyphenols (Folin–Ciocalteau Index) and
Absorbanсe (Abs) were measured according to the OIV Compendium of International Methods of Analysis of Wine
and Musts (OIV 2007). Color intensity (CI) and hue were evaluated according to the Glories method (1984). Total
anthocyanins were determined by the spectrophotometric method based on SO₂ bleaching (Ribèreau-Gayon and
Stonestreet 1965). Tannins were determined according to Ribèreau-Gayon and Stonestreet (1966). Analyses were
performed in duplicate using basic analytical equipment and a Shimadzu UV-1800 (Kyoto, Japan) UV
spectrophotometer.
Phenol was extracted from the grapes as follows: the separate extraction of berry components was carried out in
duplicate, simulating the maceration process necessary for the production of red wines (Mattivi et al., 2002; Vacca et
Briefly: berries (200 g) were cut in two with a razor blade, and seeds and skins were carefully removed from each half of the berry. The pulp on the inner face of the berry skin was removed using an end-flattened spatula in an attempt to preserve skin integrity. Skins and seeds were immediately immersed in a 200 mL solution of ethanol: water (12:88 v/v), 100 mg/L of SO₂, 5 g/L tartaric acid and a pH value adjusted to 3.2 (with NaOH) and extracted for five days at 30°C. The extracts were shaken by hand once a day. Skins and seeds were removed from the hydro-alcoholic solution after five days and the skin extract was centrifuged for 10 min at 3500 × g. Extracts were poured into dark glass bottles, flushed with nitrogen and stored at 4°C until spectrophotometric analyses.

3. Results and Discussion

3.1 Homogeneous Zones (HZ) identification after soil and geophysical mapping

In order to identify potentially different environments leading to Homogeneous Zones, we performed a standard soil mapping adapted to the specific needs of this research. The combination of a geomorphological analysis of LIDAR-derived DEM and local geological data led to the production of preliminary landscape mapping units. These mapping units (not reported) depict three different environments, namely (i) a summit landscape unit having a slope gradient of about 5-10% developed over clayey sediments with clear signs of local erosional processes, (ii) an upslope landscape unit having a slope gradient of about 25-30% developed over clayey sediments with little signs of erosional processes, (iii) a downslope landscape unit having a slope gradient of about 7-15 % developed over a colluvium landform with no signs of erosion. An EMI survey was then produced on these units to orient the pedological survey (also in terms of soil variability) and to define the boundaries of HZs.
The ECa maps obtained were used as baseline data in the pedological survey. The ECa maps (Figures 3) showed that the vineyard was clearly characterized by the presence of two major patterns of ECa, generally homogenous, corresponding to two areas: (i) summit and upper slope (red area in Fig. 3d) and (ii) downslope (bluish area in Fig. 3d). The ECa mean difference between the two areas was statistically significant (P<0.05). The difference between summit and upslope previously observed by the preliminary landscape analysis (on high resolution DTM), however, did not always correspond to the ECa mapping and, moreover, the boundaries between the two main areas identified by the ECa pattern is oblique (with respect to slope) and not linear as would be expected with DTM analysis.

Landscape mapping unit analysis was combined with ECa mapping, observing ECa homogenous and heterogeneous areas and performing 25 qualitative rapid soil observation (minipits, augers). Six soil profiles and 10 augers were localized to include major variability (Fig. 3a). The soil profiles and augers were described and sampled. Bulk and undisturbed soil samples were collected from each described soil horizon and submitted for chemical and physical analysis.

From the pedological characterization (76 soil samples in total, 51 from 6 soil profiles and 25 from augers; augers data are not shown), two main soil types were identified: Cambic Calcisol (Clayic, Aric) and Eutric Cambisol (Clayic, Aric, Colluvic) (WRB, 2014). These soils are likely to have evolved from a different parent material; the Cambisol evolved over colluvium (including traces of pumices), not present in the upper part of the vineyard, while the Calcisol evolved from the clayey sedimentary bedrock. The different origin is also expressed by the soils color: brown (10YR) for the Calcisol and yellowish (5Y) for the Cambisol.

Calcisol is richer in total carbonates than Cambisol (mean 232.7 g/kg and 41.2 g/kg respectively), with a Bk horizon at a depth of about 45 cm with the common accumulation of carbonates. This induces a differentiation of the pH between the two pedons (Calcisol mean 8.2; Cambisol mean 7.0). The organic Carbon Content (O.C) and Cation Exchange
Capacity (CEC) were higher in the Cambisol (avg. values of O.C. in the topsoil was 1.0 and 2.1% for the Calcisol and Cambisol respectively; avg. value of CEC along the soil profile was 24.3 cmol*kg⁻¹ in the Cambisol and 16.8 cmol kg⁻¹ in the Calcisol). The apparent electrical conductivity (ECa) is generally low in both soils, highlighting the absence of significant quantities of salts in solution (173 and 246 mS cm⁻¹ for the Cambisol and Calcisol respectively). The texture is clay loam in both soils.

The physical characteristics of two soil profiles, representative of the two soil types indicated, are reported in table 1. Despite the similar texture, the hydraulic properties measured in the lab showed some important differences (Tab. 1). Among them: i) Calcisol showed a pronounced vertical heterogeneity (i.e. the $K_0$ and $l$ of the Bk horizon are very different from the adjacent upper and lower horizons); ii) Cambisol showed a relative vertical homogeneity, especially in the Bw horizons; iii) despite the fact that the porosity of Calcisol is higher than that of Cambisol (see saturated soil water contents, $\theta_0$) the available water content (AWC) in the first 80 cm of soil depth was lower (i.e. Calcisol 80 mm and Cambisol 145 mm).

<table>
<thead>
<tr>
<th>Soil/HZ</th>
<th>Soil horizon and thickness (cm)</th>
<th>Particle size fraction</th>
<th>Hydrological properties</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td></td>
<td>Clay</td>
<td>Silty</td>
</tr>
<tr>
<td>Cambic</td>
<td>Ap1 0-10/20</td>
<td>31.9</td>
<td>38.1</td>
</tr>
<tr>
<td></td>
<td>Ap2 10/20-45</td>
<td>32.0</td>
<td>37.7</td>
</tr>
<tr>
<td></td>
<td>Bk 45-80</td>
<td>32.6</td>
<td>39.7</td>
</tr>
<tr>
<td></td>
<td>BC 80-105</td>
<td>33.8</td>
<td>39.3</td>
</tr>
<tr>
<td></td>
<td>CB 105-130+</td>
<td>34.9</td>
<td>37.6</td>
</tr>
</tbody>
</table>

<table>
<thead>
<tr>
<th>Soil/HZ</th>
<th>Soil horizon and thickness (cm)</th>
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<th>Hydrological properties</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td></td>
<td>Clay</td>
<td>Silty</td>
</tr>
<tr>
<td></td>
<td>Ap 0-40</td>
<td>34.2</td>
<td>31.5</td>
</tr>
<tr>
<td>Calcisol</td>
<td>Bw1 40-90</td>
<td>37.6</td>
<td>30.0</td>
</tr>
<tr>
<td></td>
<td>Bw2 90-120</td>
<td>42.9</td>
<td>29.5</td>
</tr>
<tr>
<td></td>
<td>CAM 120-160+</td>
<td>41.1</td>
<td>30.8</td>
</tr>
</tbody>
</table>

$\theta_0$: absent; $K_0$: few fine sub-rounded pumiceous stones

Integrating the soil and geophysical survey with the physical, hydrological and chemical soil analysis made it possible to separate two main Homogeneous Zones: CAL: Cambic Calcisol (Clayic, Aric) (Fig. 4) developing at the summit and in the uplope landscape position and CAM: Eutric Cambisol (Clayic, Aric, Colluvic) (Fig. 4) developing in the downslope landscape position.
3.2 Modelling application (potential CWSI estimation)

The potential Crop water stress index (CWSI) data were obtained by analyzing the water balance in the soil–vegetation–atmosphere (SVA) system in the two HZs using the SWAP hydrological model applied over eleven years of daily climate data (2003-2013). This information is especially important because it helps evaluate the dynamics of the water stress of the soils present in the study area, representing a very powerful tool for vineyard planning.

In particular, the level of CWSI between the two HZs in each plant phenological phase (Fig. 2 and 5) can be compared, defining the functional Homogeneous Zones (fHZs) on the basis of their similarity in functional behavior in relation to vines and correlate it to different plant responses in terms of must quality and plant production.

On average, in CAL the potential CWSI was two times higher (13.9%) than CAM (5.9%), with clearly different increases during the cropping season (from flowering to harvesting) (Tab. 2, Fig. 2 and 5).

The maximum values of CWSI were obtained during berry formation with average values of 32.7% (+- 20.7) for CAL and 13.5% (+-9.6) for CAM.

From the results of potential CWSI analysis, it is clear that the two representative soils of the study area, under the same climate and plant conditions, show a different susceptibility to vineyard water stress. These results mean that the two identified CAL and CAM HZs behave as Homogeneous Zones also in terms of their functional behavior. To describe this output of the microzoning procedure from here onwards we shall refer to CAL and CAM functional Homogeneous Zones (fHZs) as CAL and CAM fHZs.

The statistical ANOVA analysis on CWSI over the seasons showed a significant difference, with an alpha of 0.02 between the two soils. During the different phenological phases only the Berry ripening phase showed a significant differences with an alpha of 0.02. This behavior can be explained considering that both soils started the growing season with an optimum water content (accumulated during the winter), but during the season, the reduction of rainfall, the increase of ET, and the effect of plant water uptake emphasize the physical differences of these two SPA systems. Concerning the high variability of SD during berry ripening, the differences depend on different climate conditions.
conditions, in particular rainfall amount, during the 11 years analyzed. We can clearly identify two very dry years (2003 and 2007) and two very wet years (2005 and 2010).

During the simulated years, the average rainfall in the period from 18 August to 15 October was 105 mm with an SD of 46.9 (44% variation).

Fig. 5. The average values and standard deviation of the potential Crop Water Stress Index (CWSI) during the cropping season (reported in terms of phenological phases) of vines cultivated in the representative soils of homogeneous zones (CAL and CAM). The simulations were performed with SWAP models during the vintages from 2003 to 2013.

3.3 GIS analysis

Each functional Homogeneous Zone was then analyzed with respect to the variability of environmental characteristics derived from the high resolution DMS and DTM (Tab. 2; Fig. 6). The elevation and slope of CAL was higher compared to those of CAM; their mean difference is always significant ($P<0.05$) and it was about 15 m for the elevation and 4.7 % for the slope gradient. These differences are consistent with both geomorphic and soil settings.

The aspect shows a Nord-West and West orientation respectively for CAL and CAM fHZs. This aspect difference ($P<0.05$) induces - in the vineyard rows - a differentiation in terms of total potential insolation ($P<0.05$) during the cropping season (1 April to 15 October) which is about 55 kwh/m² higher in the case of the CAM fHZ. Such differences are mainly due to the direct vineyard row insolation. This different insolation was tested directly on the Penman-Monteith equation; results showed a negligible effect on evapotranspiration and then on CWSI, as predicted by the model. Probably a direct effect on grape bunches in terms of temperature could be realizable but this aspect was not treated in this work.
3.4 Vineyard records (crop/must measurements)

The experimental plots were identified inside the two fHZs and the phenological and physiological vine data were collected on 27 plants (randomly sampled at each measurement date) over three years (2011-2013).

Even though the experimental plots had the same cultivar (Aglianico), the same rootstocks (1103P) and the same management, the crop responses in terms of biomass development and must quality were very different. Plants of CAM showed more vigour when compared to those of CAL (at fruit thinning, an average value of 11.1 bunches/plant versus 10.0 were measured, with a peak of 14.6 versus 8.7 bunches/plant in the year 2011). Despite a very similar number of bunches/plant at harvesting (average value of 4.6 bunches/plant for CAM and CAL respectively), at harvest time the plant production of CAM was generally higher (1.81 ± 0.29 kg/plant) compared to that of CAL (0.97 ± 0.36 kg/plant). The last results regarding the different berry weight and volume recognized during the three years (Tab. 2) are in agreement.

The analyses carried out on grape bunches over the three years of measurement showed a very robust qualitative differentiation between the two fHZs. Investigated parameters like sugar, anthocyanins, polyphenols in the skin, colour intensity, tannins in the skin and pH were always higher in CAL during berry ripening if compared to CAM. However, the titratable acidity, bunch volume and weight of 100 berries were lower in the CAL. The first results of microvinification showed higher values of ethanol (12.2 %v for CAM and 13.3 %v for CAL), colour intensity (7.8 for CAM and 12.8 for CAL) and tannins (2.9 g/l for CAM and 4.6 g/l for CAL) in CAL than in CAM.

ANOVA analysis showed that during the three years only the total anthocyanins and color intensity of the grape must were significantly affected by soil characteristics (p<0.05). These results are in agreement with previous findings showing that moderate, and not severe, water stress or drought stress increases anthocyanins concentrations in berry skins (Ojeda et al., 2002). Among the three years considered, 2013 stands out because the grapes showed a lower...
content of sugars, a higher weight and a lower density of berries, mainly due to the fact that during 2013 the temperatures were lower and the ripening season was rainy. As a consequence, degree-day accumulation was slower and the berries larger due to their watering.

In both fHZs the minimum absolute values of LWP (-0.37 and -0.40 MPa in CAM and in CAL, respectively) were registered at the beginning of crop season, while the maximum absolute values were reached at the end of August (-1.65 and -1.85 MPa in CAM and in CAL, respectively). Nevertheless, during the three years of measurements the plants belonging to Calcisol (CAL fHZ, up-slope) faced a more intense water stress than those in Cambisol (CAM fHZ, down-slope), during the whole season. Consequently, stomatal conductance values during the three years agreed with those of LWP, with the Calcisol plants experiencing lower values, and thus lower transpiration rates than those of Cambisol. Assimilation rates followed the same behavior of stomatal conductance, highlighting that in CAM the plants had a more pronounced photosynthetic activity than those of CAL; in agreement with that, also the quantum yield of photosystem PSII in leaves adapted to light (PHIPSII) showed that CAM plants were more efficient than CAL plants in capturing the energy of the light absorbed by the photosystem PSII. Both photosynthetic activity and PHI PSII responded proportionally to the different chlorophyll a content, with the CAM plants showing the highest values. The Leaf Area index (LAI) was lower in CAL plants (avg. 1.28) than in CAM plants (avg. 1.48), as a consequence of the more severe water stress suffered by plants grown in the former. Moreover, during the three years of the experiment (2011-2013), the differences between the two fHZs of all parameters monitored on the plant were significant, with p<0.001 (T-test, Two-tailed). Only the LAI showed a P value of 0.014.
### Tab. 2. Summary of results obtained in CAL and CAM HZs and CAL and CAM fHZs: (i) GIS analysis on DSM and DTM; (ii) Simulation modelling application; (iii) Plant monitoring and the characteristics of bunches during the three years of monitoring (2011 to 2013) on Aglianico cv.

<table>
<thead>
<tr>
<th>Environmental characteristics and plant responses</th>
<th>CAL HZ / CAL fHZ</th>
<th>CAM HZ / CAM fHZ</th>
</tr>
</thead>
<tbody>
<tr>
<td><strong>Elevation (m.s.l.)</strong></td>
<td>363.8 (± 6.5)</td>
<td>348.5 (± 6.5)</td>
</tr>
<tr>
<td><strong>Slope (%)</strong></td>
<td>23.2 (± 6.9)</td>
<td>18.5 (± 5.9)</td>
</tr>
<tr>
<td><strong>Aspect (°N)</strong></td>
<td>300.8 (± 20.0)</td>
<td>276.6 (± 26.2)</td>
</tr>
<tr>
<td><strong>TWI</strong></td>
<td>-2.3 (± 1.2)</td>
<td>-2.4 (± 1.1)</td>
</tr>
<tr>
<td><strong>Pot. direct insolation (kWh/m²)</strong></td>
<td>635.9 (± 294.8)</td>
<td>708.1 (± 326.3)</td>
</tr>
<tr>
<td><strong>Pot. diffuse insolation (kWh/m²)</strong></td>
<td>136.6 (± 20.9)</td>
<td>137.3 (± 21.0)</td>
</tr>
<tr>
<td><strong>Tot. Pot. insolation (kWh/m²)</strong></td>
<td>786.9 (± 502.7)</td>
<td>845.4 (± 335.7)</td>
</tr>
<tr>
<td><strong>Shoot growth</strong></td>
<td>1.0 % (± 1.6)</td>
<td>0.23% (± 0.0)</td>
</tr>
<tr>
<td><strong>Flowering</strong></td>
<td>6.9 % (± 9.9)</td>
<td>1.4 % (± 10.6)</td>
</tr>
<tr>
<td><strong>Berry Formation</strong></td>
<td>10.8 % (± 8.8)</td>
<td>13.5 % (± 9.6)</td>
</tr>
<tr>
<td><strong>Berry ripening</strong></td>
<td>32.7 % (± 20.7)</td>
<td>13.5 % (± 9.6)</td>
</tr>
<tr>
<td><strong>CWSI over cropping season</strong></td>
<td>13.9 % (± 8.6)</td>
<td>5.9 % (± 3.9)</td>
</tr>
<tr>
<td><strong>Leaf Water Potential (LWP) (Mpa)</strong></td>
<td>Mean -1.12 (± 0.33)</td>
<td>-0.92 (± 0.27)</td>
</tr>
<tr>
<td><strong>Min (abs.)</strong></td>
<td>-0.4</td>
<td>-0.37</td>
</tr>
<tr>
<td><strong>Max (abs.)</strong></td>
<td>-1.85</td>
<td>-1.65</td>
</tr>
<tr>
<td><strong>Chlorophyll content of leaves (CCL)</strong></td>
<td>13.9 (± 4.3)</td>
<td>18.6 (± 6.67)</td>
</tr>
<tr>
<td><strong>Phot. CO₂ Assimilation (µmol m⁻² s⁻¹)</strong></td>
<td>10.3 (± 6.0)</td>
<td>15.7 (± 5.9)</td>
</tr>
<tr>
<td><strong>Stomatal conductance (mol m⁻² s⁻¹)</strong></td>
<td>0.15 (± 0.1)</td>
<td>0.23 (± 0.1)</td>
</tr>
<tr>
<td><strong>Instantaneous Water Use Efficiency (WUE) (µmol mol⁻¹)</strong></td>
<td>Mean 81.6 (± 34.0)</td>
<td>72.3 (± 24.5)</td>
</tr>
<tr>
<td><strong>Leaf Area Index (m²/m²)</strong></td>
<td>1.28 (± 0.20)</td>
<td>1.48 (± 0.40)</td>
</tr>
<tr>
<td><strong>Effective quantum yield of photosystem PSI</strong></td>
<td>Mean 0.11 (± 0.04)</td>
<td>0.15 (± 0.04)</td>
</tr>
</tbody>
</table>

* The potential insolation refers to the period 1 April to 15 October

** Average values calculated over 11 years (2003-2013)

*** Average over three seasons of measurements (2011 to 2013)

### Discussion

The effects of soil combination and climate on vine responses in terms of must characteristics and wine quality are well reported in literature as the basis of the terroir concept. With this work, conducted in a small study area, the usefulness of the adopted microzoning procedure was tested by quantifying the effect of soil properties on plant responses and must characteristics.
The effectiveness of the results achieved lies in the specific experimental setup that was conducted in the same geomorphic land system, under the same climate conditions (only 90 m of slope, with about 15 m difference in elevation) on the same plant cultivar (Aglianico monoclonal population) and under the same vine management.

We believe that under these conditions, the large-scale soil survey supported by the EMI survey represented a very good cost/benefit approach to investigating vineyard potentialities in view of a soil-plant-climate relationship study. Indeed the EMI approach made it possible to immediately identify a homogeneous pattern on which to focus the sampling activities.

The identification and mapping at vineyard scale of the main soil types included in the functional Homogeneous Zones, CAL fHZ, Cambic Calcisol (Clayic, Aric), and CAM fHZ, Eutric Cambisol (Clayic, Aric, Colluvic), allowed us to estimate and study the different behaviour, in terms of crop water stress, of these two soils in the soil-plant-climate system. This discriminated their different abilities in affecting the quality of a wine.

The potential CWSI referred to the last eleven years was very different between the two HZs. The variability expressed by this index for each phenological phases was very high. This was due to the high variability of the weather data used as input for the simulation modelling application.

These important differences between the soil-crop-climate behaviour are mainly due to the different hydrological behaviour, because of the very different hydraulic properties of the soil (see Table 1).

It is essential to notice that the use of simulation models like the one used for this study requires an accurate phase of model calibration and preferably the use of measured data, rather than those estimated by methods (such as pedofunctions) that usually tend to smooth out the soil hydrological properties between otherwise very different soils. This is also the case if the AWC is used as an indicator for crop water availability: the use of PFT could produce a clear mistake, also if the studied soils present similar texture, as reported in section 3.1.

This was our case, where the two soils fell under the same textural class while having very different hydrological behavior. Hence, the measurement of the hydrological properties was of primary importance in differentiating the two environments.

The behaviour of each soil-plant and atmosphere system (CAL and CAM) was investigated in depth, making the conclusion about their different ability to produce crop water stress very solid. Thus, it was possible to identify two functional Homogeneous Zones, corresponding to the previous HZs.

The CAL fHZ represents a system where the Aglianico is subjected to a water stress that is twice as strong as that occurring in the Cambisol fHZ, with a progressive differentiation from flowering to harvesting.

This is clearly in agreement with the water stress felt by plants during the three years of monitoring (avg. 22% of LWP increase in the CAL). In addition, the r Pearson of CWSI estimated by the model and by the LWP measured on-field was 0.98. Moreover, the different behaviour described by the potential CWSI was also confirmed by other plant physiological measurements including: (i) stomatal conductance (the plants in CAL experienced lower values, and thus lower transpiration rates, than those of CAM), (ii) assimilation of CO2 rates (in CAM the plants had a more pronounced photosynthetic activity than those of CAL) and (iii) the quantum yield of photosystem PSII in leaves adapted to light (PHIPSII) (plants in CAM were more efficient than CAL plants at capturing the light energy absorbed by the photosysterm PSII).

From the enological viewpoint, grapes analyzed in this study showed important differences. In CAL fHZ, the grapes were richer in sugars, anthocyanins, total polyphenols and had a lower content of total acids. Considering that Aglianico wines, as traditionally produced, are generally rather acid, astringent and easily downfall red colour (Gambuti et al., 2007), then these data clearly suggest that grapes belonging to CAL fHZ can produce wine with a more balanced taste that is more alcoholic and less acidic.

Grapes from CAM fHZ showed a higher extractable polyphenol content than those from CAM fHZ, indicating that a more aged wine can be obtained from this part of the vineyard. On the basis of grape sugar content, wines obtained from this fHZ should be also characterized by a higher content of ethanol (13.5% v/V with respect to 12% v/V for CAM fHZ wines) and should show a more intense colour because of the content of native pigments (anthocyanins) extracted from skins and their colour intensity. In contrast, CAM fHZ berries showed a lower content of total anthocyanins extracted from skins and a similar content of total polyphenols and tannins extracted from seeds. Taking into account the facts that: i) anthocyanins are mainly extracted during the first phases of red vinification (consisting in the maceration of whole berries during must fermentation), ii) complete extraction from the seeds requires the
berry skins and seeds to be in contact for a longer time with must-wine, and iii) seed tannins are more astringent than
skin tannins (Gambuti et al., 2006), these data suggest that a specific winemaking procedure, such as short
maceration, could help obtain from CAM fHZ grapes a red wine with a good colour intensity, which is not astringent
and which is easier to drink.
Therefore, the enological potentials of grapes belonging to the two sites are very different. By applying the proper
winemaking procedure it is possible to obtain a more ready-to-drink wine from the CAM fHZ site and a long ageing
wine from the CAL fHZ site.
In conclusion, the use of a model output is a useful approach to evaluating and comparing the effects of the CWS in
vines induced by soils.
Anyhow, important prerequisites that should be considered are: i) model calibration (if previous data are available to
calibrate it), ii) preferably measured data, iii) same plant cultivar and iv) same climatic and plant management
conditions. This was our case, built up to investigate and compare the “soil suitability to grape production”, limiting
the effects of other environmental variables. This comparison would not have been feasible with different cultivars
(each of which responds differently to water stress), or different boundary conditions.
Moreover, our approach could be improved towards the precision farming considering the soil spatial variability inside
of each fHZs in the modelling application, thereby providing also an information about the uncertainty of model
predictions. But, the used of latter approach is strictly limited by the availability of soil information spatially distributed
needed to apply the simulation model (verticals and horizontals information, e.g. soil horizons depth and thickness and
their hydraulic properties).
However, the great potential of the dynamic simulation models applied in this context can be seen. In fact, once the
characteristics and parameters of the different SPA systems (i.e. LAI, climate data, cultivar relation between CWSI and
quality must parameters, etc.) are known, it is potentially possible to estimate, as done in this work, what could be the
plant responses to water stress and therefore its effects in must and wine quality, in any soil-plant combination or in
any boundary conditions, including plant responses to future climate changes. This could have imaginable positive
effects on future land use and management planning, for instance the choice of the most suitable plant varieties for
specific production targets or the opportunity to apply drip irrigation systems in order to control the plant water
status with the aim of improving quality or maintaining the current level. This concept is very similar to the approach
reported in literature for other crops between the yield response and water stress (or water deficit) (see Menenti et
al., 2014, or Monaco et al., 2014).
Finally, the GIS analysis of high resolution DTM and DSM showed that differences in terms of slope and elevation
between the two identified fHZs, at vineyard scale, were low and not very important. On the other hand, the aspect
and the potential insolation calculated over the cropping season (1 April to 15 October) showed that in the CAM fHZ
the plants receive 7% or more total potential insolation during the cropping season compared to those cultivated in
the CAL fHZ. This condition strengthens our results and confirms the hypothesis that for the scale of our work, soil
drives the Aglianico plant expression in terms of must quality and then wine quality.

5 Conclusions
The procedures adopted for viticulture microzoning which include (i) standard large-scale soil mapping, (ii) geophysical
mapping and (iii) soil-plant water stress evaluation on the identified fHZs, have shown their robustness in terms of
their effects on plant, grape, must and wine quality.
The inclusion of the soil-plant water stress evaluation was fundamental because plant water status affects the
characteristics of the grape must, skin and seeds of the Aglianico vine.
In particular the study has shown: i) the importance of the hydropedology approach in knowing the soil properties, in
order to come to a complete characterization of the different pedo-environments (also recognizable at field scale)
aimed at viticultural zoning, equal to the well-recognized importance for soil chemical properties; ii) the link between
must characteristics and soil characteristics, particularly CWSI estimated using a simulation model. This can be
considered as preliminary information for zoning and planning the vineyard plant (e.g without having local data for
calibrating and validating the model); iii) the need to transform the soil map into a functional map in the viticultural
zoning procedures, where the soils are evaluated dynamically on the basis of soil-plant and atmosphere system
behaviour (e.g. soil water balance), with the definition of functional homogeneous zones (fHZs) for the vine; and (iv)
the potentiality of this approach to explore future prospects in terms of more effective grape variety selection and
precision irrigation application to overcome the high CWSI values expected from climate change.

Acknowledgements

We acknowledge Dr. A. Erbaggio, P. Caputo, A. Delle Cave for the field measurements and Mrs Orefice for soil
hydraulic properties measurements.

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